



## GEOFYSIK

Regionala flygminningar saknas, men omfattande elektromagnetiska markmätningar (slingram) har utförts över delar av Kvilleberggrundens. Mätningar finns från strax söder om Sjöläja till norra kultigränsen för 22E NV, över ett område kring Björkvatnet (22E, 3-4a) samt över prospekteringsobjekten Lilljället, Småvattenbranna (St. Jörn) och Kvartjärmen (5-6d). Mätningarna lämnar i regel värdefull geologisk information, eftersom den ökade ledningsförmågan hos de grafitiska fylltorna och skifferna ger upphov till utmåliga anomaler.

## MALMER OCH INDUSTRIELLA MINERAL

Ur malmgeologisk synpunkt har de lagertornigt uppträckta (stratiforma) kismalmerna tilldragit sig det största intresset sedan de första upptäckterna gjordes i början av 1900-talet. De klassas numera som exhalativ sedimentära, d.v.s. avsatta i samband med vulkanisk verksamhet genom lösningsar som utfällt på havsbotten eller i de närmast underliggande ånnu ej konsoliderade bergarterna. Malmerna domineras i regel av kompaktartad svavelkis med varierande halter av zink, koppar och i regel obetydligt bly samt visst innehåll av silver och guld (se tabell på kartans framsida). Impregneringar artade mineraliseringar förekommer också.

**Ankarvattnetmalmen** har undersökts av SGU (31 st borrhål) och beräknas innehålla 753 000 ton malm med halter enl. tabell. **Jörnlemmalmen** ägs f.n. av Boliden Mineral AB, 103 borrhål har slagits och tonnaget uppskattas till 612 000 ton (halter, se tabell). **Björkvattnetmalmen** har uppskattats av SGU (30 st borrhål). Totalt tonnaget uppgår till endast 132 000 ton med relativt låga mettalhalten (se tabell). **Jörpajietti** (Ola Andersson) har likaledes undersökts av SGU (5 st borrhål) men är tydlig ifråga om vissa kismalmsanläggningar som finns listade i tabellen eller inlagda i kartan.

Några parter av de ultrafamiska kropparna har utmållats på grund av förhöjda halter av *krom* (i minskat kromit), och förutsättningarna för nickel-/koboltutvinning har undersökts, framför allt under 1970-talet. Ingendera av metallerna är dock med nuvarande förutsättningar ekonomiskt utvinningsbar, ej heller mineralen *magnesiit* eller *talk*.

**Glimmer**, främst för användning inom den elektriska industrien, bröts under avspärrningsåren 1941–1945 i ett flertal små fyndigheter, 15–25 km NNE om Gädde. Förekomsten utgörs delvis av pegmatitgångar, delvis av lagerformatade kroppar och linser av grovpladig glimmer i glimmerskiffer. Det ungefäriga läget av ett antal brutna fyndigheter har inlagts på kartorna och listats i tabellen.

## DESCRIPTION

### General geology

The bedrock within the four map sheets (22D-22E Tjäderöarna) forms part of the Scandinavian Caledonides. The rocks are late Proterozoic – early Paleozoic in age, although the high-grade rocks in the east (Seve) may contain older Precambrian elements. Most Caledonian rocks in Scandinavia, and all the units in the present area, are allochthonous and have been thrust east- or southeastwards onto the Baltoscandian platform. Regionally, the Caledonides are divided in ascending order into the Autochthon and the Lower, Middle, Upper and Uppermost Allochthon (Jürgen 1972, Gee et al. 1985). All tectonostratigraphic units in the area belong to the Upper Allochthon, which is composed of higher-grade Seve rocks, metamorphosed under amphibolite-facies conditions, locally with high-P assemblages (eclogites), and overlying Koli nappe composed of lower-grade, greenschist – lower amphibolite facies rocks.

The Seve rocks are dominated by quartz-rich metasedimentary units which were probably deposited along the western edge of the late Proterozoic – early Paleozoic continent Baltic, and metabasic rocks which represent intrusions and extrusions related to the opening of the Palaeozoic Iapetus Ocean (Prestwich 1851; Röhl 1972). The Koli nappe is the oldest allochthonous unit in the area, and the Lower, Middle and upper Koli (Gee et al. 1985). The Lower Koli is interpreted to represent ensimatic arc-basin complexes that probably developed closer to the ‘European’ side of the Iapetus, whereas the Middle and Upper Koli nappes represent more exotic ensimatic arc-basin sequences which probably developed along the Laurentian (‘American’) side of the Iapetus (Stephens and Gee 1985, Stephens 1988). All the Seve and Koli units have a complex tectonic and metamorphic history (Dallmeyer and Gee 1988). The Seve units were affected by a Late Cambrian – Early Ordovician event, locally producing high-pressure assemblages, and all units were affected by Ordovician-Silurian deformation and metamorphism. The various complexes were successively brought together along the suture zone formed during collision of the continental Baltic and Laurentia in the Late Silurian – Early Devonian. At this time the Caledonian activity faded out.

### Tectonostratigraphic units

The geological units distinguished on the maps are principally lithological or lithostatigraphic in character. The map legend is identical for the four sheets although all rock types are not present on each individual map sheet. The rock sequence regionally forms a tectonostratigraphy where the different, generally NW-dipping units are separated by major and minor thrusts, as demonstrated by the structural inset map and its legend.

### Seve units

About two thirds of the bedrock within the map area are composed by Seve rocks. It should be noted that several of the eastern units, dominated by quartzite or meta-arkose with mafic rock intercalations, were classified by Strömborg (1984) as part of the Middle Allochthon. The state of metamorphism, including the presence of eclogites and retro-eclogites in several units, the structural pattern, and regional comparisons and correlations argue for inclusion in the Upper Allochthon.

The **Sljuten** unit (Bakker 1978) is used here in an extended sense as a name for several smaller units dominated by quartzite or meta-arkose with mafic rock intercalations; the largest of which is called the Fiskäfjället Amphibolite. The amphibolites are often garnetiferous, sometimes porphyritic with feldspar megacrysts; cross-cutting dolerites occur. On the 22E Frostviken map sheets, the Gakkeljälet and Sjöpilen Formations are present only in the SE quadrangle.

The **Gakkeljälet** unit (Van Roermund 1976) is used here in an extended sense as a name for several units dominated by quartzite or meta-arkose with mafic rock intercalations; the largest of which is called the Gakkeljälet Amphibolite. The amphibolites are often garnetiferous, sometimes porphyritic with feldspar megacrysts; cross-cutting dolerites occur. On the 22E Frostviken map sheets, the Gakkeljälet and Sjöpilen Formations are present only in the SE quadrangle.

The **Björkvattnet** unit is a grouping of several amphibolite-dominated complexes, including the Björkvattnet Formation (Sjöstrand 1978). They enclose the Eremaja Formation (Van Roermund 1976), the Blériot ‘Enheit’ (Blériot 1978), the Sjöläja Amphibolite (Sjöstrand 1978) and the Sipmik-Värtan ‘Formation’ (Kardoe 1978); overlying are the Giper Amphibolite (Winter 1974) and the Grutenvalle (Kardoe 1978) and Tjökkola (Van Roermund 1977) ‘Formations’. The amphibolites are locally garnet-bearing and metasedimentary intercalations of garnet-biotite-muscovite schist, quartzite and marble also occur.

The **Ertske Lens** occurs as a tectonic lens or a detached, recumbent, isoclinal fold-hinge within the above-mentioned amphibolites. It is composed of two lithologically distinct subunits. The **Lejaren** unit including the Lejaren ‘Formation’ (Sjöstrand 1978), the Krönike Quartzite (Winter 1974) and the Rieksvarto ‘Formation’ (Kardoe 1978) are high-grade quartz-rich metasedimentary, often foliated and topographic index mineral-rich quartzite facies, often with garnetiferous intercalations. The **Avardo** unit (acc. to Sjöstrand 1978) is a kyanite-sillimanite-K-feldspar gneiss which forms the host rock to most of the eclogites in the map area. Peak metamorphic conditions here have been estimated to be 18.0±1.0 kb and 780±50°C (Van Roermund 1985).

The **Blåsjöjälen** unit is a grouping of several amphibolite-dominated complexes, including the Blåsjöjälen Formation (Sjöstrand 1978). They enclose the Blåsjön-Björnhöjden area, Caledonides of W. Jämtland, Sweden. – Unpubl. map, 1:50 000, Univ. Leiden.

The **Gakkeljälet** unit (Van Roermund 1976) is used here in an extended sense as a name for several smaller units dominated by quartzite or meta-arkose with mafic rock intercalations; the largest of which is called the Gakkeljälet Amphibolite. The amphibolites are often garnetiferous, sometimes porphyritic with feldspar megacrysts; cross-cutting dolerites occur. On the 22E Frostviken map sheets, the Gakkeljälet and Sjöpilen Formations are present only in the SE quadrangle.

The **Blåsjön** unit (Van Roermund 1976) is used here in an extended sense as a name for several small units dominated by quartzite or meta-arkose with mafic rock intercalations; the largest of which is called the Gakkeljälet Amphibolite. The amphibolites are often garnetiferous, sometimes porphyritic with feldspar megacrysts; cross-cutting dolerites occur. On the 22E Frostviken map sheets, the Gakkeljälet and Sjöpilen Formations are present only in the SE quadrangle.

The **Lejaren** unit including the Lejaren ‘Formation’ (Sjöstrand 1978), the Krönike Quartzite (Winter 1974) and the Rieksvarto ‘Formation’ (Kardoe 1978) are high-grade quartz-rich metasedimentary, often foliated and topographic index mineral-rich quartzite facies, often with garnetiferous intercalations. The **Avardo** unit (acc. to Sjöstrand 1978) is a kyanite-sillimanite-K-feldspar gneiss which forms the host rock to most of the eclogites in the map area. Peak metamorphic conditions here have been estimated to be 18.0±1.0 kb and 780±50°C (Van Roermund 1985).

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sequence, the eclogite-bearing Ertske Lens is placed at a structurally lower level, beneath the Central Belt, probably intercalated within the Eastern Schist and Amphibolite Belt of Trouw (1973).

### Transition units

This term was used by Sjöstrand (1978) in describing various tectonic units of uncertain affinity between the Seve and the overlying Koli rocks. Without detailed microscopical investigations it is often difficult to decide in the field whether the rocks represent the lower prograde part of the Koli, or retrograded to an earlier stage of metamorphism. The rocks are deposited in the platformal and megasequence areas, as opposed to the platformal and megasequence areas which are part of the late Precambrian - Silurian sequence of the Autoclinion and Lower Allochthon, deposited on the Baltoscandian platform. Only Lower Koli and Middle Koli rocks are present within the area.

### Koli units

The greenschist facies metamorphic rocks of the Upper Allochthon are referred to as Koli. They represent the Cambro-Silurian sedimentary and volcanic eugeochemical assemblages, deposited in exotic westerly areas, as opposed to the platformal and megasequence areas which are part of the late Precambrian - Silurian sequence of the Autoclinion and Lower Allochthon, deposited on the Baltoscandian platform. Only Lower Koli and Middle Koli rocks are present within the area.

Lower Koli rocks form a continuous, up to 5 km wide zone diagonally through map sheets. Definition of the character and limits of this unit is well-established further north in its type area around Björkvattnet (Sjöstrand 1978). The Stikke Nappe has been tentatively traced through the northern part of the area, and the Björkvattnet Nappe has been demonstrated only in the Frostviken map sheets. A U/Pb zircon dating (Clæssøn et al. 1982) has defined an age of 488±5 Ma (Arénig?), which is consistent with the inferred stratigraphic age. The quartzite/marble horizon above the volcanic rocks has been correlated with the Vojtová/Städäla Formation of Ashgillian age in the type area.

In the map area, the tectonic contact at the top of the Björkvattnet Nappe has been demonstrated only between the Lower Koli and the Stora Blåsjön (Sundblad 1981). Further northeast and southwest, the thrust has been tentatively traced slightly above the above-mentioned quartzite/marble horizon, locally including some calcareous phyllites (Garbenschiefer) which may represent the Lovfjäll Phyllite (Kulling 1993). The presence of black phyllites (equivalents to the Broken Formation) and associated mafic volcanic rocks and intrusions makes this zone susceptible to tectonic dislocations.

Middle Koli is represented by two different tectonic units in the present area. The **Stikke Nappe** derives its name from the Stekenjkk area (Stephens 1982) where the felsic-dominated Stekenjkk Volcanics (Stephens 1982) are exposed. The Stikke Nappe is the uppermost tectonic unit in the area, and through the 22D-E map sheets, where, south of western Kvambergsvatnet, it has been named the Skogsbäcken Volcanics (Sjöstrand 1978). The stratigraphic sequence of the Stikke Nappe is inverted. Thus, the Basalt-Quartz-Korrbekong Formation (Nilsson 1964) is structurally overlain, but stratigraphically underlain, by variable, dark, often granitic phyllites and mafic volcanites (Remdalen Group of Zachrisson 1969) and stratigraphically overlain by the underlying calcareous phyllites (Blåsjö Phyllite of Nilsson 1964). U/Pb zircon dating (Clæssøn et al. 1982) gives a minimum age of 476±1 Ma for the Skogsbäcken Volcanics and an age of 440±2 Ma for felsic, trondjemite intrusions in the (stratigraphically) lower part of the Blåsjö Phyllite. Thus, the age of the rocks in the Stikke Nappe is probably Ordovician.

The uppermost tectonic unit of the area, along the national border to Norway, is represented by the Middle Koli **Lepikvattnet Nappe** (Zachrisson 1969). The thrust at its base is geologically significant around lake Lepikvattnet. Although difficult to pin-point in the present map sheets, it has to be traced somewhere within the sequence of dark, often granitic phyllites. A calcareous phyllite, the Bräkfall Phyllite (Nilsson 1964), forms the major formation of the nappe. Characteristic rock types in the Lepikvattnet area are polymictic conglomerates, coarse fragment-bearing metagreywackes and the Bjurås limestone. Some of these units impinge on the northernmost edge of 22E NW.

## LITERATURE

SGU = Sveriges geologiska undersökning  
GFF = Geologiska Föreningen i Stockholm Förhandling

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